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Abstract

The origin of Carlin-type or sediment-hosted, disseminated gold deposits of the Great Basin, the major source of gold in the United States, is poorly understood. We propose that Eocene magmatism was the heat source that drove the hydrothermal systems that generated these deposits in the Carlin trend and Independence Mountains in northern Nevada. This interpretation is based on a strong spatial and temporal association of Eocene intrusive-volcanic centers with the gold deposits of this region. Our new work and published $^{40}\text{Ar}/^{39}\text{Ar}$ dates indicate that magmatism was particularly intense between 39 and 40 Ma throughout northeastern Nevada, especially in and around the area of gold deposits. Carlin-type deposits may have formed preferentially during Eocene magmatism because it was (1) more intense in the area than other magmatic episodes, (2) somehow compositionally distinct, or (3) accompanied by extension that promoted hydrothermal flow. However, large-scale extension does not appear to have been a factor in generating Carlin-type deposits.

First Page Preview

Eocene magmatism: The heat source for Carlin-type gold deposits of northern Nevada

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ABSTRACT

The origin of Carlin-type or sediment-hosted, disseminated gold deposits of the Great Basin, the major source of gold in the United States, is poorly understood. We propose that Eocene magmatism was the heat source that drove the hydrothermal systems that generated these deposits in the Carlin trend and Independence Mountains in northern Nevada. This interpretation is based on a strong spatial and temporal association of Eocene intrusive-volcanic centers with the gold deposits of this region. Our new work and published $^{40}\text{Ar}/^{39}\text{Ar}$ dates indicate that magmatism was particularly intense between 39 and 40 Ma throughout northeastern Nevada, especially in and around the area of gold deposits. Carlin-type deposits may have formed preferentially during Eocene magmatism because it was (1) more intense in the area than other magmatic episodes, (2) somehow compositionally distinct, or (3) accompanied by extension that promoted hydrothermal flow. However, large-scale extension does not appear to have been a factor in generating Carlin-type deposits.

INTRODUCTION

Carlin-type deposits in the Great Basin have been intensely studied, yet their origin remains controversial (Sillitoe and Bonham, 1990; Archart et al., 1993; Kuehn and Rose, 1995; Thorman et al., 1995; Ilchik and Barton, 1997; Oppliger et al., 1997). Controversy centers on the heat source that drove hydrothermal flow and the age of the deposits. Some interpret igneous intrusions to be the heat source (Sillitoe and Bonham, 1990; Archart et al., 1993; Hofstra, 1995; Thorman et al., 1995; Leonardson and Rahn, 1996; Groff et al., 1997). For example, Sillitoe and Bonham interpret Carlin-type deposits to be distal to porphyry intrusions. Others interpret the heat source to be elevated crustal heat flow related to extension and not to magmatism (Kuehn and Rose, 1995; Ilchik and Barton, 1997); Ilchik and Barton (1997, p. 273) claimed "Carlin-type deposits lack demonstrable time-space links to intrusive centers." Still others consider the heat source to be coupled magmatism and extension (Seedorff, 1991; Oppliger et al., 1997).

The age of Carlin-type deposits is debated because they are difficult to date directly. Host rocks are mostly Paleozoic sedimentary rocks that have undergone multiple episodes of diagenesis, alteration, and metamorphism. The resulting fine-grained alteration minerals are neither easily dated nor clearly tied to gold deposition. Nevertheless, recent work indicates that many formed in the Eocene (Hofstra, 1995; Emsbo et al., 1996; Leonardson and Rahn, 1996; Phinisey et al., 1996; Rota, 1996; Groff et al., 1997). For example, at the giant Goldstrike deposit in the Carlin trend (Fig. 1), a dacitic dike dated as 39.3 ± 0.4 Ma ($^{40}\text{Ar}/^{39}\text{Ar}$; biotite) was initially interpreted to be postmineralization (Archart et al., 1993). Emsbo et al. (1996) found that the dike had locally undergone the same alteration and had the same geochemical enrichment as ore-bearing rocks and concluded that the dike predated ore. Leonardson and Rahn (1996) concluded that this dike was emplaced approximately at the end of mineralization. In the Independence Mountains (Fig. 1), a basaltic dike dated as 40.8 ± 0.1 Ma ($^{40}\text{Ar}/^{39}\text{Ar}$; whole rock) is altered and locally contains ore; a weakly propylitized quartz monzonite dike dated as 39.2 ± 0.1 Ma ($^{40}\text{Ar}/^{39}\text{Ar}$; hornblende) is outside the main area of alteration and may postdate mineralization (Phinisey et al., 1996). Phinisey and others concluded that ore was deposited contemporaneously with or closely following Eocene magmatism.

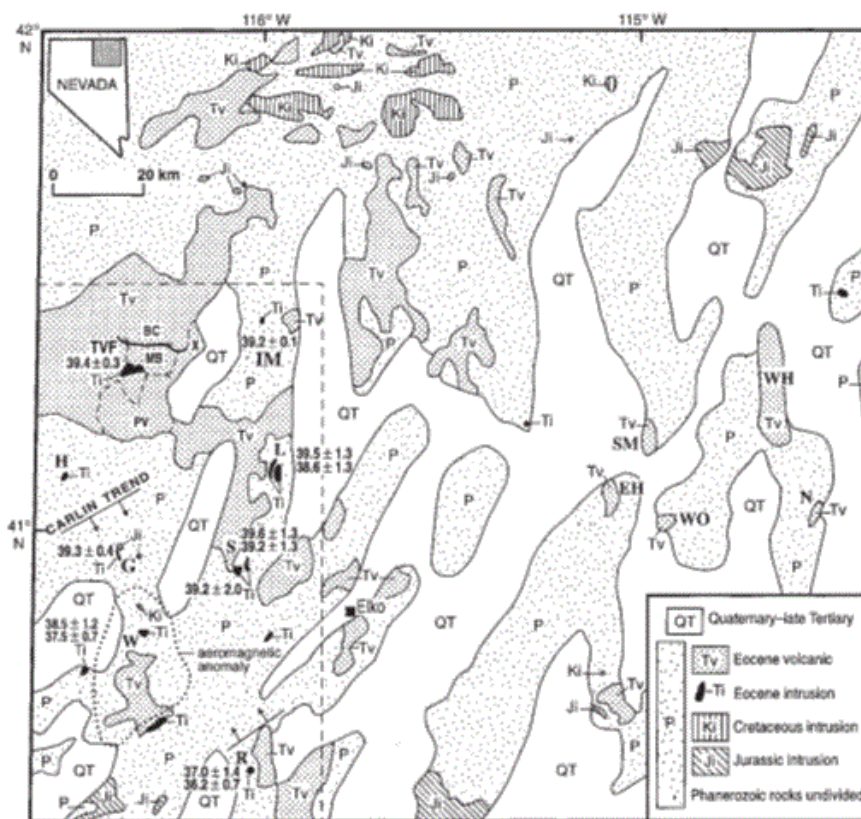


Figure 1. Generalized geologic map of northeastern Nevada showing major gold areas, Carlin trend (G—Goldstrike deposit) and Independence Mountains (IM), in relation to Eocene igneous rocks. Gold deposits are concentrated in area of most intense Eocene magmatism and mostly away from areas of Cretaceous or Jurassic intrusions. Aeromagnetic anomaly (from Hildenbrand and Kucks, 1988) along southwestern edge of Carlin trend indicates large intrusion, probably of Eocene age. Magnetic anomalies are also associated with each of Eocene igneous centers in area but are not shown. TVF—Tuscarora volcanic field (PV, Pleasant Valley volcanic complex; MB, Mount Blitzen volcanic center; BC, Big Cottonwood Canyon caldera; X, Sixmile Canyon lavas); H, Hatter stock; L, Lone Mountain; S, Swales Mountain; W, Welches Canyon; R, Railroad district; EH, East Humboldt Range; SM, Snake Mountains; WO, Wood Hills; WH, Windermere Hills; N, Nanny Creek. Dashed box outlines subarea of Table 2.

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